

factors that regulate the growth of phytoplankton. Longer day length and higher light intensity have been shown to be associated with higher biomass production than shorter days and lower light intensities; moreover, complete utilization of nitrate and phosphate occurs in a shorter amount of time under conditions involving longer day lengths and higher light intensities (Meseck *et al.*, 2005).

The northwestern plateau is a unique tectonic feature. The bedrock within its catchment consists of sedimentary and metamorphic rocks, as well as small amounts of igneous rocks. Because of the high altitude (1000–2000 m), the vegetation in the catchment areas of most of the lakes consists mainly of alpine meadows and alpine steppes. The vegetation on the plateau includes predominantly C3 plants and varies from swampy meadow vegetation to mountain shrubs; C4 plants play only a minor role. The southwestern plateau is the middle tier of the eastern slope of the Himalayas. It is a complex geomorphological unit that is characterized by low latitudes and high altitudes (1000–2000 m), as well as large altitude gradients. This area is affected by both the southeastern (Pacific Ocean) and southwestern (Indian Ocean) monsoons. The main vegetation is subtropical evergreen broad-leaved forest, which co-exists or is mixed with deciduous broad-leaved forest. Superimposed on the natural climatic changes are more recent anthropogenic influences that have been caused by industrialization and changes in agricultural practices. It is important to differentiate between anthropogenic impacts and those caused by natural environmental changes.

This project provides a comprehensive study of the composition, preservation and origin of bulk OM and lipids in lacustrine sediments from six lakes. The lakes include northwestern and southwestern plateau lakes, which cover a range of latitudes extending from 25°47' (N) to 41°58' (N). Natural factors that may influence the composition of sedimentary OM, such as light intensity, MAWT, MAP, pH, and TDS, differ among the lakes. To identify the spatial distribution and changes in the source of sedimentary OM, along with the factors that influence the different plateau lakes, TOC and TN concentrations, $\delta^{13}\text{C}_{\text{org}}$ and $\delta^{15}\text{N}_{\text{total}}$ values, and the compositions of biomarkers (AHs, *n*-alkanols and FAs) in the surface sediments were analysed.

METHODS

Lake selection and sediment collection

Six lakes (Lake Hongfeng, Lake Chenghai, Lake Erhai, Lake Bosten, Lake Sugan and Lake Qinghai) were selected from the southwestern and northwestern plateaus in China. The parameters MAWT, MAP, pH and TDS clearly differ among these study lakes, due to their widely

varying latitudes (which range from 25°47' (N) to 41°58' (N)) and the effects of human activities. Four sediment cores were collected from the centre of each lake in 2014 using a piston-percussion corer fitted with 58 mm i.d. Perspex tubes. The cores were sectioned into 1-cm intervals and freeze-dried immediately after collection. To increase the representativeness of the studied samples, the surface sections (0–1 cm) of all four sediment cores from each lake were combined for analysis.

Elemental and stable isotope analyses

Subsamples for elemental (TOC and TN) and bulk stable isotope composition analyses were acidified with 1 M HCl before analysis to remove carbonates. The concentrations of TOC and TN, as well as the $\text{C}_{\text{org}}/\text{N}_{\text{total}}$ (atomic) ratios, were determined using a CHNS Vario E1 III elemental analyser. Carbon and nitrogen isotope analyses were conducted on a Thermo Finnigan Delta Plus XL mass spectrometer connected to a Flash EA 1112 elemental analyser via a Finnigan MAT ConFlo III interface. The stable isotopic ratios $\delta^{13}\text{C}_{\text{org}}$ and $\delta^{15}\text{N}_{\text{total}}$ are reported relative to the Vienna Pee Dee Belemnite and air standards, respectively. The instrumental analytical precision was 0.1‰ for $\delta^{13}\text{C}_{\text{org}}$ and 0.2‰ for $\delta^{15}\text{N}_{\text{total}}$.

Extraction and fractionation of free and bound lipids

The samples were Soxhlet extracted for 72 h with $\text{CH}_2\text{Cl}_2/\text{MeOH}$ (9:1 v/v). Sulfur was removed through the addition of activated Cu. The free lipids were fractionated using silica gel-alumina column chromatography (50 mm \times 5 mm i.d.). The saturated hydrocarbons were eluted with 20 mL of hexane, followed by the free alkanols and free fatty acid (FFA) fractions, which were successively eluted using 20 mL of 20% ethyl acetate (EtOAc) in hexane and 20 mL of methanol (MeOH). The extracts were saponified with 0.5 M KOH in MeOH under reflux for 2 h to release the ester-bound lipids. The mixtures were centrifuged, and the supernatant was decanted. The bound alkanols were extracted with *n*-hexane/ Et_2O (9:1 v/v). After acidification to pH 1 through the addition of HCl, the bound fatty acids (BFAs) were extracted with CH_2Cl_2 . All fractions were analysed using gas chromatography (GC) and GC mass spectrometry (GC-MS). Prior to the GC and GC-MS analyses, the free and bound alkanol fractions were derivatized with acetic anhydride/pyridine (1:1 v/v). The FFA and BFA fractions were methylated with MeOH (100°C, 1 h) to convert them to FA methyl esters.

Gas chromatography and gas chromatography-mass spectrometry

GC was conducted using a Finnigan Trace GC instrument equipped with an HP-5 fused silica column (50 m \times 0.32 mm \times 0.25 μm film thickness). The carrier gas used

was N_2 , and the oven temperature program was $70^\circ C$ (2 min) to $290^\circ C$ (held 30 min) at $3^\circ C \text{ min}^{-1}$. The GC-MS analyses were performed using a Finnigan Platform II mass spectrometer coupled to a Hewlett-Packard (HP) 6890 GC instrument and an HP-5 fused silica column ($50 \text{ m} \times 0.32 \text{ mm} \times 0.25 \mu\text{m}$ film thickness). The temperature program was $70^\circ C$ (5 min) to $290^\circ C$ (held 30 min) at $3^\circ C \text{ min}^{-1}$. The *n*-alkanes, alkanols, and FAs were quantified through integration of the peak areas in total ion current (TIC) plots. Perdeuterated eicosane was used as an internal standard, and the response factor for individual AHs, alkanols, and FAs relative to the standard was assumed to be 1.0.

RESULTS

Bulk composition of organic matter

The TOC contents varied between 2.0 and 4.9% for the sedimentary samples. The TOC contents were generally higher in the sediments obtained from the southwestern

plateau lakes than those obtained from the northwestern plateau lakes in China (Fig. 1). The $C_{\text{org}}/N_{\text{total}}$ values ranged between 8.2 and 11.4 for the sedimentary OM. The $\delta^{13}C_{\text{org}}$ values for most of the lakes varied between -29.4‰ and -24.0‰ . The only exception was the sediment from Lake Sugan, which had $\delta^{13}C_{\text{org}}$ values of -18.8‰ . In this study, the $\delta^{15}N_{\text{total}}$ values ranged from $+4.6$ to $+8.7\text{‰}$.

Composition and distribution of free and bound lipids

Nineteen AHs, 15 *n*-alkanols and 32 FAs were identified. The FAs included normal saturated FAs ($C_{12:0}$ to $C_{28:0}$), branched FAs (e.g., *iso*- $C_{15:0}$ and *anteiso*- $C_{15:0}$), and monounsaturated FAs (e.g., $C_{16:1}$ and $C_{18:1}$). The compounds were divided into several subgroups according to their terrestrial, algal, bacterial, and mixed origins.

Aliphatic hydrocarbons

The AHs in the sedimentary samples were dominated by *n*-alkanes ranging from C_{15} to C_{31} . As shown in Fig. 2,

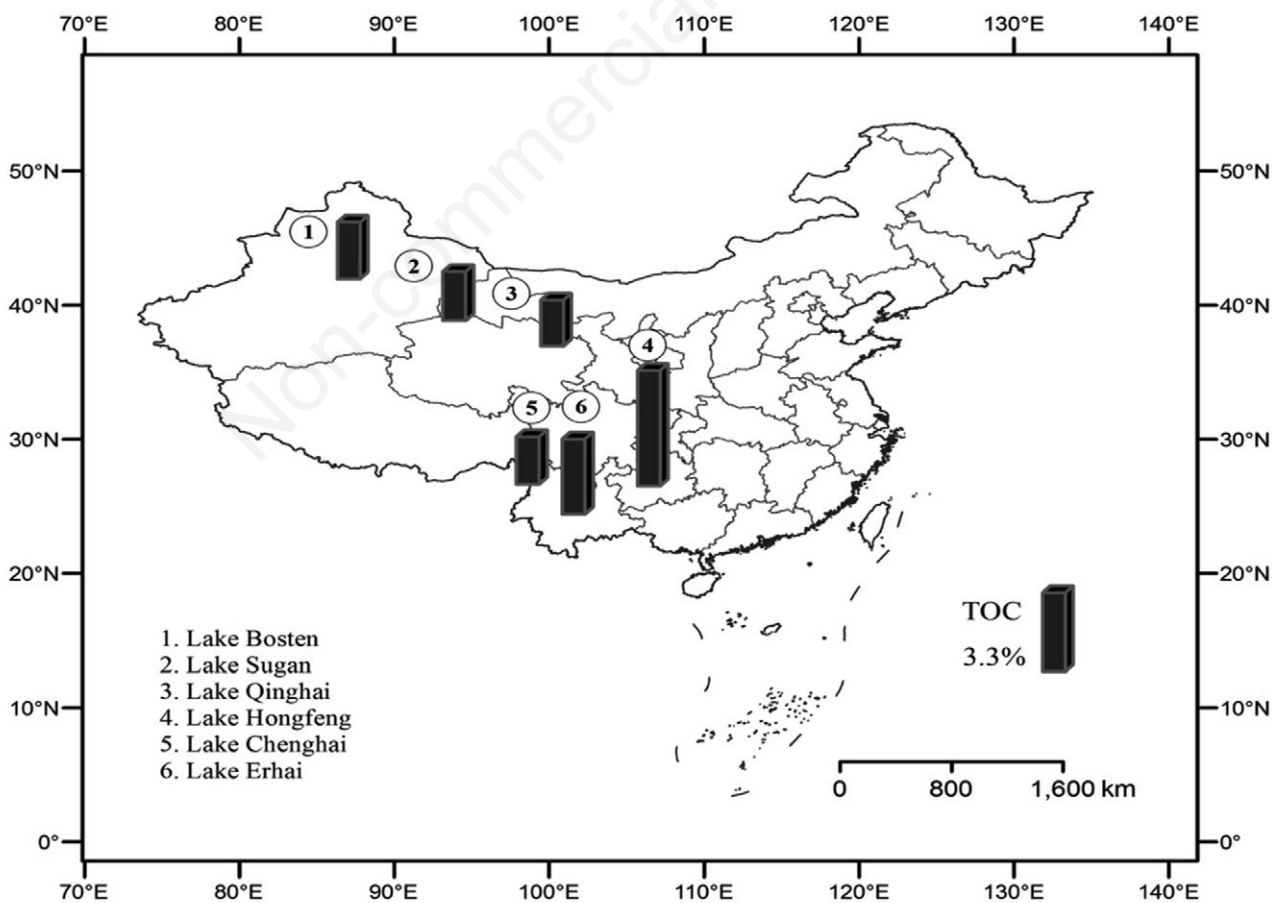


Fig. 1. Sketch map of PR China showing the locations and TOC distributions in the surface sediments from northwestern and southwestern plateau lakes in China.

the distributions of *n*-alkanes differed considerably among the lakes studied. The *n*-alkanes found in the sediments obtained from the southwestern plateau lakes were generally dominated by plankton-derived C₁₅-C₂₀ components with a maximum around C₁₇ or C₁₈. In contrast, the C₂₅-C₃₁ *n*-alkanes derived from terrestrial higher plants peaking at C₂₇, C₂₉, or C₃₁ dominated the sedimentary *n*-alkanes from the northwestern plateau lakes. Pristane (Pr) and phytane (Ph), which are products of the geological alteration of phytol and other natural isoprenoid products, were found in all samples.

Fatty acids

FAs were the most abundant lipids, and they made up 80.1-91.1% of the total lipids (TL) by concentration. TL represents the sum of all lipid compounds measured, including AHs, alkanols, and FAs. The concentrations of BFAs ranged from 3569.8 to 10,776.4 µg gOC₋₁, and the FFA concentrations ranged from 80.9 to 590.0 µg gOC₋₁. The FA composition showed that the saturated fatty acids (SFAs) were dominated by C_{14:0}, C_{16:0}, and C_{18:0} in the sediment samples. Except for the BFAs in the sediment from

Lake Hongfeng, which display a maximum at *i*-C_{15:0} (accounting for 39.3% of total BFA), the TICs for the FFAs and BFAs in all other samples were characterized by a unimodal distribution pattern, and the short-chain *n*-C_{16:0} component was predominant (Fig. 3); *n*-C_{16:0} constituted 24.9-40.6% and 19.0-30.1% of the FFAs and BFAs, respectively. *Iso*-acids in the range C₁₂ to C₁₈ occurred in the BFAs and those from C₁₃ to C₁₇ occurred in the FFAs. On the other hand, *anteiso*-acids ranging from C₁₃ to C₁₇ were found in the BFAs, and only *anteiso*-C₁₅ was found to occur in the FFAs. Polyunsaturated fatty acids (PUFAs) are normally associated with algae. However, the PUFA concentrations were below detectable levels in all of our samples.

Aliphatic alkanols

The *n*-alkanol fractions extracted from the lake sediments had carbon chain lengths that ranged from C₁₄ to C₂₈, and a strong predominance of even over odd chains was noted (Fig. 4). The concentration of free *n*-alkanols ranged from 58.0 to 492.7 µg gOC₋₁ (1.0-8.2% of TLs), whereas the bound *n*-alkanols varied from 142.0 to 382.6

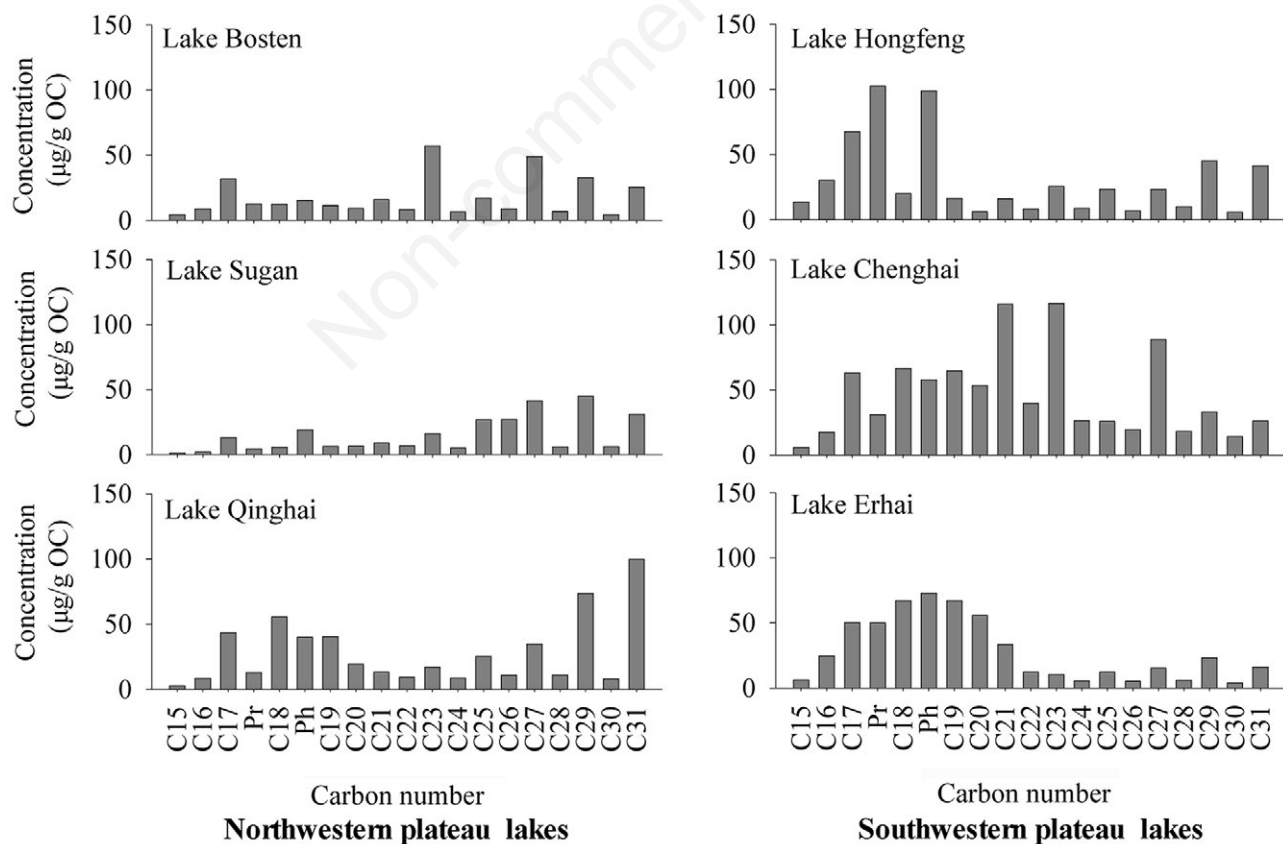


Fig. 2. Distribution of aliphatic hydrocarbons in the surface sediments.

$\mu\text{g gOC}_{-1}$ (2.6–8.6% of TLs). There were marked differences between the distributions of free and bound *n*-alkanols. For example, the bound *n*-alkanols were dominated by short-chain compounds (C_{14} – C_{22}), and many samples were not found to contain any long-chain compounds (C_{24} – C_{28}). However, in some of the sediment samples, the free *n*-alkanols displayed a bimodal distribution with a maximum at C_{18} or C_{22} , a secondary maximum at C_{26} or C_{28} (Fig. 4) and a strong predominance of even-chain lengths.

DISCUSSION

Sources of organic matter

The parameters $\text{C}_{\text{org}}/\text{N}_{\text{total}}$, $\delta^{13}\text{C}_{\text{org}}$ and $\delta^{15}\text{N}_{\text{total}}$, which are useful bulk indicators of the sources of organic matter examined in paleolimnological reconstructions, can be used to assess the relative contribution of allochthonous *versus* autochthonous components. The ratio of $\text{C}_{\text{org}}/\text{N}_{\text{total}}$ usually ranges from 5 to 8 in plankton and bacteria and from 20 to 100 in terrestrial organic matter (Jasper and Gagosian, 1989). Tab. 1 shows that the $\text{C}_{\text{org}}/\text{N}_{\text{total}}$ ratios range from 8.2 to 11.4 and are intermediate between values that are characteristic of autochthonous and allochthonous inputs of organic matter, indicating a mixed source

for the organic matter found in the sediment samples. The $\text{C}_{\text{org}}/\text{N}_{\text{total}}$ ratios are closer to the range associated with typical autochthonous sources than to that associated with allochthonous sources. This result indicates that the sedimentary OM from the northwestern and southwestern plateau lakes in China is derived mainly from aquatic algae and bacteria. The $\delta^{13}\text{C}_{\text{org}}$ values of most of the sediment samples obtained from the studied lakes varied between -29.4‰ and -24.0‰ (Tab. 1), and these values are characteristic of both lacustrine algae and C_3 land plants (Meyers, 2003). The only exception was the sediment from Lake Sugan, which had $\delta^{13}\text{C}_{\text{org}}$ values of -18.8‰ . The heavier $\delta^{13}\text{C}_{\text{org}}$ values from Lake Sugan may have been caused by the higher TDS values noted at that lake (Tab. 1). The $\delta^{13}\text{C}$ values of OM from aquatic sources in the lake respond to the carbon isotopic composition of the lake water dissolved inorganic carbon (DIC) source, and high $\delta^{13}\text{C}_{\text{DIC}}$ values are associated with increasing TDS values (Li *et al.*, 2012).

Terrigenous detrital OM is generally characterized by a low $\delta^{15}\text{N}_{\text{total}}$ ($+0.5\text{‰}$) signature, whereas aquatic components have a relatively high $\delta^{15}\text{N}_{\text{total}}$ ($+8.5\text{‰}$) value (Thornton and McManus, 1994). In addition, $\delta^{15}\text{N}_{\text{total}}$ values can also be used to investigate the sources and mechanisms of pollution. For example, typical $\delta^{15}\text{N}_{\text{total}}$ values range from -2 to $+4\text{‰}$ for commercial fertilizers, $+3$ to

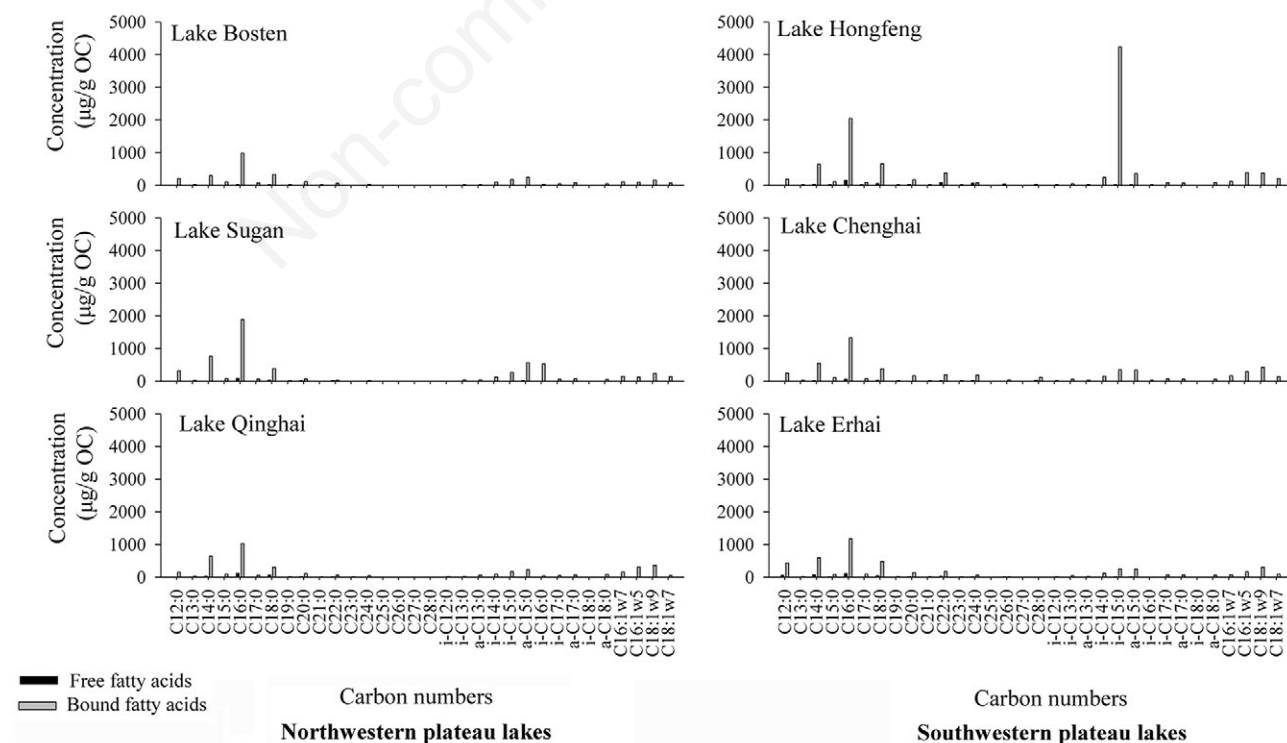


Fig. 3. Distribution of fatty acids in the surface sediments.

+8‰ for soil organic nitrogen nitrate and +10 to +20‰ for human and animal waste nitrate (Aravena *et al.*, 1993). In this study, the $\delta^{15}\text{N}_{\text{total}}$ values ranged from +4.6 to +8.7‰ (Tab. 1). These values provide further evidence for a mixed source of OM, and they indicate that the inputs from algae and bacteria were dominant in these surface sediment samples. A few common ratios, such as TAR_{HC} ($\text{TAR}_{\text{HC}} = (C_{27} + C_{29} + C_{31}) / (C_{15} + C_{17} + C_{19})$) n-alkanes, TAR_{AK} ($\text{TAR}_{\text{AK}} = (C_{24} + C_{26} + C_{28}) / (C_{14} + C_{16} + C_{18})$) n-alkanols, and TAR_{FA} ($\text{TAR}_{\text{FA}} = (n\text{-}C_{24:0} + n\text{-}C_{26:0} +$

$n\text{-}C_{28:0}) / (n\text{-}C_{14:0} + n\text{-}C_{16:0} + n\text{-}C_{18:0})$) fatty acids (Meyers, 2003; Fang *et al.*, 2014), served as molecular geochemical proxies for land-derived *versus* algal contributions of lipids in this study. Tab. 2 displays the values of TAR_{HC} , TAR_{AK} , and TAR_{FA} for the OM in the sedimentary samples. The n-alkanes are less susceptible to degradation than other components of organic matter in sediments; thus, the n-alkane distributions are more useful for identifying sources of organic matter. Terrigenous OM is commonly richer in n-alkanes than algal and bacteria material,

Tab. 1. Basic data and total organic matter indexes for six plateau lakes in China.

Lake	Sampling date	Lake area* (km ²)	MAWT* (°C)	Sampling site depth (m)	Catchment area* (km ²)	TDS* (g L ⁻¹)	MAP* (mm)	pH*	TOC (%)	TN (%)	C _{org} /N _{total}	$\delta^{13}\text{C}_{\text{org}}$	$\delta^{15}\text{N}_{\text{total}}$
Lake Bosten	9-Sep	992.0	7.9	16.5	22000	1.9	62.3	8.6	2.4	0.3	9.2	-26.9	4.6
Lake Sugan	13-Sep	108.0	2.7	6.1	7000	31.8	77.6	8.5	2.1	0.3	8.2	-18.8	5.3
Lake Qinghai	16-Sep	4340.0	5.8	28.2	29661	15.5	336.6	9.2	2.0	0.3	8.9	-28.1	6.0
Lake Hongfeng	21-Sep	57.2	14.1	45.0	1596	0.2	1200.6	8.5	4.9	0.5	11.4	-29.4	8.7
Lake Chenghai	25-Sep	77.2	18.7	36.9	318	1.0	725.5	9.2	2.0	0.3	9.2	-26.5	5.2
Lake Erhai	28-Sep	249.0	15.3	21.5	2470	0.3	1100.4	8.4	3.2	0.4	8.8	-26.8	4.7

MAWT, mean annual surface water temperatures; TDS, total dissolved solids; MAP, mean annual precipitation; TOC, total organic carbon; TN, total nitrogen; *obtained from the China Lake Scientific Database and published data.

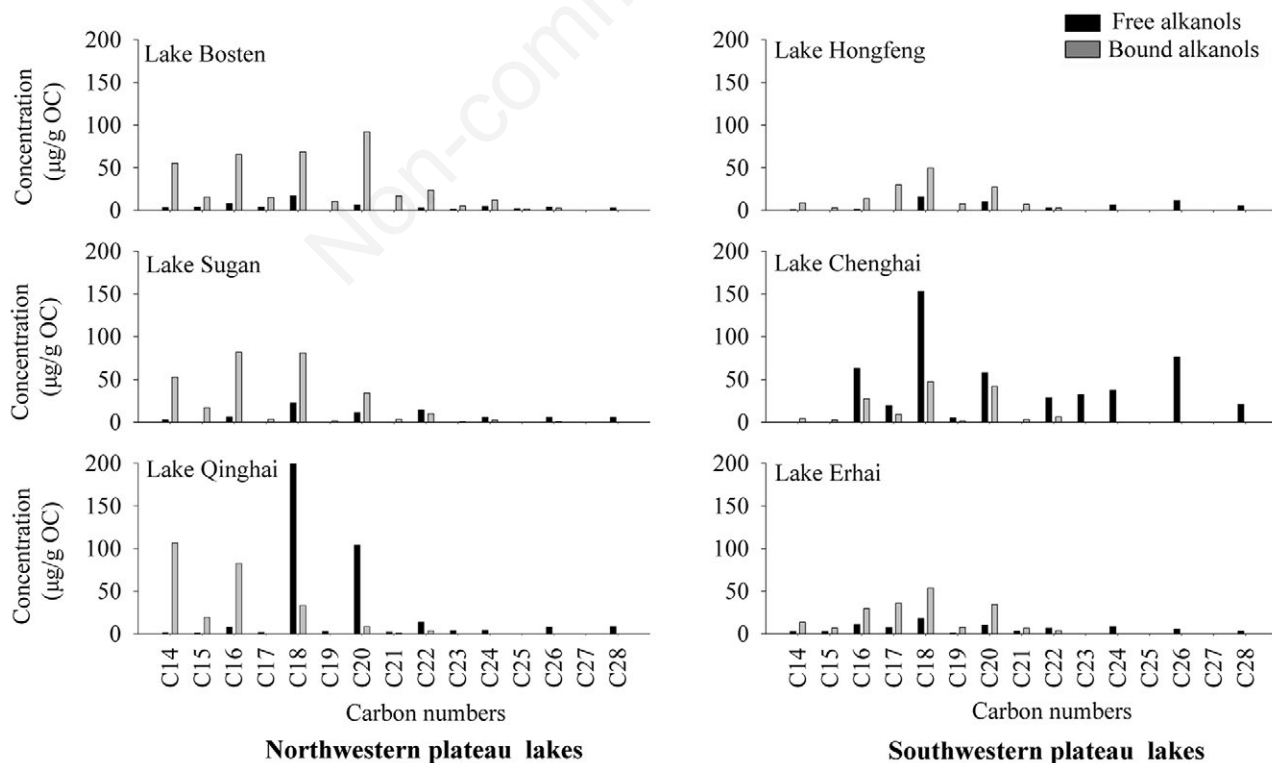


Fig. 4. Distribution of n-alkanols in the surface sediments.

and TAR_{HC} values are therefore useful in identifying changes in the proportions of terrigenous vs aquatic contributions of hydrocarbons. TAR_{HC} ratios are sensitive to contributions to sedimentary OM from vascular plants. Therefore, changes in watershed vegetation that lead to increased contributions from tree and shrub litter cause TAR_{HC} values to rise significantly (Silliman *et al.*, 1996). The TAR_{HC} ratios showed a decreasing trend from the high-latitude to low-latitude lakes (Tab. 2). This pattern is probably due to the enhanced aquatic production that occurs in lakes at lower latitudes, which causes high aquatic hydrocarbon concentrations and dilutes the terrigenous signal. Low TAR_{HC} values indicate a higher proportion of aquatic OM in sediments. More aquatic OM was present in the sediment collected from the southwestern plateau lakes (Tab. 2). These elevated levels of aquatic OM were mainly due to the higher nutrient salt levels, light intensities, and MAWTs in these lakes, which can promote algal and bacterial growth and reproduction (Meseck *et al.*, 2005; Cao *et al.*, 2011). In addition, factors that influence *n*-alkane distributions in lacustrine sediments included ongoing anthropogenic hydrocarbon inputs, such as industrial discharges (from petroleum distribution and refining), shipping activities (dry docking), and sewage outfalls. Secondary *n*-alkanes derived from the decomposition of terrestrial hydrocarbons can also affect *n*-alkane distributions in lacustrine sediments (Meseck *et al.*, 2005; Lü and Zhai, 2006).

The ratios of biomarker FAs can help to summarize the sedimentary FA record. These source identifiers were used to calculate the ratios of terrigenous-to-aquatic FAs, $TAR_{FA} = (n-C_{24:0} + n-C_{26:0} + n-C_{28:0}) / (n-C_{14:0} + n-C_{16:0} + n-C_{18:0})$. High TAR_{FA} values can indicate relatively abundant lipids from terrigenous sources relative to aquatic sources, but they can also indicate degradation of aquatic FAs relative to land-derived components (Tenzer *et al.*, 1999).

Selective degradation and other diagenetic effects commonly overprint FA source signatures. Short-chain acids are often preferentially degraded by microbes during early diagenesis (Ho and Meyers, 1994). This type of preferential degradation increases TAR_{FA} values; however, no such elevated values were evident in these sedimentary samples. Microbial synthesis of secondary FAs from primary OM produces shorter-chain components (Kawamura *et al.*, 1987), which can depress TAR_{FA} values. All of the TAR_{FA} values of the BFAs and FFAs were low (Tab. 2). These low values indicate that either algal contributions were predominant in the FAs, or microbial reworking of sedimentary OM overprinted the original source characteristics of the FA components. Both land-plant and algal FAs would need to have been heavily reworked to yield the low TAR_{FA} values found in these sediments.

The *n*-alkanol compositions of algae and bacteria are dominated by C_{16} to C_{22} *n*-alkanols (Volkman *et al.*, 1999). C_{22} and C_{24} *n*-alkanols are mainly derived from submerged and floating plants (Ficken *et al.*, 1998), and C_{22} to C_{28} *n*-alkanols have been found in land and emergent water plants. The *n*-alkanol terrigenous-to-aquatic ratio is defined as $TAR_{AK} = (C_{24} + C_{26} + C_{28}) / (C_{14} + C_{16} + C_{18})$. The TAR_{AK} values for the sedimentary samples ranged from 0.1 to 1.3, and these values ranged from 0.0 to 0.1 for free and bound *n*-alkanols, respectively (Tab. 2). The relatively small contributions of the more readily preserved long-chain *n*-FAs and alkanols indicate that the initial contributions of land-plant lipids to sedimentary samples were smaller compared with the algal and bacterial contributions. Unlike the general dominance of vascular plant-wax contributions seen in the TAR_{HC} values from the northwestern plateau lakes, the TAR_{FA} and TAR_{AK} values showed that, in general, FAs and *n*-alkanols from aquatic sources were predominant. Since *n*-alkanes are more resistant to microbial degradation than *n*-alkanols and FAs

Tab. 2. Biomarker parameters for aliphatic hydrocarbons (AHs), *n*-alkanols, and fatty acids (FAs).

Lake	Aliphatic hydrocarbons					Alkanols				Fatty acids							
	TAR_{HC}	CPI_{15-20}	CPI_{25-31}	$C_{16ratio}$	Pr/Ph	CPI_{16-28}	TAR_{AK}		CPI_{16-28}		TAR_{FA}		BRFA /SSFA		MUFA /SSFA		MUFA /BRFA
							Free	Bound	Free	Bound	Free	Bound	Free	Bound	Bound	Bound	
Lake Bosten	2.3	1.6	5.1	22.9	0.8	5.3	4.8	0.4	0.1	8.5	8.5	0.4	0.0	0.1	0.3	0.2	0.5
Lake Sugan	5.7	1.5	2.9	128.0	0.2	n. a.	20.0	0.5	0.0	11.6	16.2	0.1	0.0	0.2	0.5	0.2	0.4
Lake Qinghai	2.4	1.0	5.7	36.1	0.3	32.0	131.0	0.1	0.0	17.9	10.0	0.1	0.0	0.1	0.3	0.4	1.0
Lake Hongfeng	1.1	1.7	4.4	12.2	0.9	n. a.	1.9	1.3	0.0	8.8	17.1	0.6	0.0	0.2	1.3	0.3	0.2
Lake Chenghai	1.1	1.0	2.8	45.7	0.7	7.0	8.2	0.6	0.0	10.2	10.5	0.5	0.2	0.2	0.4	0.4	0.8
Lake Erhai	0.4	0.8	3.5	16.8	0.5	5.0	2.1	0.5	0.0	12.1	9.2	0.2	0.0	0.1	0.3	0.2	0.7

n.a., not available. $C_{16ratio} = \sum n\text{-alkanes} / n\text{-}C_{16}$; $CPI_{15-20} = (C_{15} + C_{17} + C_{19}) / (C_{16} + C_{18} + C_{20})$; $CPI_{25-31} = (C_{25} + C_{27} + C_{29} + C_{31}) / (C_{26} + 2C_{28} + C_{30})$; $CPI_{16-28} = 1/2[(C_{16} + 2C_{18} + 2C_{20} + 2C_{22} + 2C_{24} + 2C_{26} + C_{28}) / (C_{17} + C_{19} + C_{21} + C_{23} + C_{25} + C_{27})]$; $TAR_{HC} = (C_{27} + C_{29} + C_{31}) / (C_{15} + C_{17} + C_{19})$; $TAR_{AK} = (C_{24} + C_{26} + C_{28}) / (C_{14} + C_{16} + C_{18})$; $TAR_{FA} = (n\text{-}C_{24:0} + n\text{-}C_{26:0} + n\text{-}C_{28:0}) / (n\text{-}C_{14:0} + n\text{-}C_{16:0} + n\text{-}C_{18:0})$; $BRFA/SSFA = \sum(BRFA) / \sum(SSFA)$; $MUFA/SSFA = \sum(MUFA) / \sum(SSFA)$; $MUFA/BRFA = \sum(MUFA) / \sum(BRFA)$.

(Zhou *et al.*, 2010), we attribute the discrepancies among the *n*-alkane, *n*-alkanol, and FA source patterns in the sedimentary OM to post-depositional alteration of the primary alkanol and FA distributions and at least partial replacement by secondary microbial alkanol and FA biomarkers (Zhou *et al.*, 2010). The short-chain *n*-alkane distributions displayed no significant preference for odd or even numbers of carbons (see the CPI₁₅₋₂₀ values shown in Tab. 2) in most of the sedimentary samples. They may be derived from the membrane lipids of microorganisms, fossil fuel inputs, or microbial reworking of plant *n*-alkanes. The carbon preference index (CPI₂₅₋₃₁) is an indicator of the *n*-alkane source. Hydrocarbons composed of a mixture of compounds originating from land plants show a predominance of odd-numbered carbon chains with CPI₂₅₋₃₁ within the range of approximately 5-10 (Jeng, 2006). Higher CPI₂₅₋₃₁ values found in sediment indicate a greater contribution from vascular plants (Rieley *et al.*, 1991; Hedges and Prahl, 1993), and CPI₂₅₋₃₁ values close to unity may indicate inputs from microorganisms, recycled OM, and/or petroleum (Kennicutt *et al.*, 1987). The *n*-alkane CPI₂₅₋₃₁ values in the sedimentary samples were found to vary between 2.9 and 5.7, and they showed an overall decreasing trend from the high-latitude to low-latitude lakes (see CPI₂₅₋₃₁ values; Tab. 2). This result indicates that warmer conditions favour greater microbial activity, and it can also be attributed to inputs of petrogenic hydrocarbons from human activities.

As *n*-C₁₆ is rarely found in biolipids, the C₁₆ ratio, which is defined as the sum of all *n*-alkanes/*n*-C₁₆, is usually high (≈ 50) in biogenic materials compared with the relatively low values (≈ 15) observed in petroleum-contaminated samples (Colombo *et al.*, 1989). In the present study, the C₁₆ ratios ranged from 12.2 to 128.0. Lake Erhai and Lake Hongfeng have likely been contaminated by petrogenic hydrocarbons, based on their CPI₂₅₋₃₁ values < 3 , C₁₆ ratios < 15 and lower CPI₁₅₋₂₀ values, which are consistent with oil-contaminated sediments (Tab. 2). The lakes that have probably been contaminated by petroleum hydrocarbons are located mainly on the southwestern plateau, where they are subject to influences from anthropogenic activities, leading to larger amounts of industrial effluent and domestic sewage entering these lakes (Wang *et al.*, 2007). In uncontaminated sediments, the Pr/Ph ratio is higher than 1, typically between 3 and 5 (Steinhauer and Boehm, 1992). In addition, a Pr/Ph ratio less than 1 represents anoxic conditions, while a value greater than 1 reflects oxic conditions (Didyk *et al.*, 1978). The Pr/Ph ratios were less than 1 in all of the sedimentary samples examined in this study (Tab. 2), which suggests that the surface sediments are subject to anoxic conditions and/or have received material from petroleum sources.

The characteristic FAs can be used to calculate the relative dominance of different sources in lake sediments.

The proportions of the short-chain saturated fatty acids (SSFAs, *n*-C_{12:0}-C_{20:0}), MUFAs and BRFAs to the total aquatic BFAs vary from 38.8 to 67.3%, 7.1 to 36.5%, and 10.0 to 50.5%, respectively, whereas the proportions of the SSFAs and BRFAs to the total aquatic FFAs are in the range of 78.5 to 94.1% and 5.9 to 21.5%. Rajendran *et al.* (1997) found that MUFAs were predominantly present in aerobic bacteria (60.5%), whereas anaerobic bacteria and sulfate-reducing bacteria (SRB) contained approximately equal amounts of BRFAs (74.0% and 78.1%). In addition, Gram-positive bacteria are also reported to contain high amounts of BRFAs. The predominance of BRFAs and MUFAs in different groups of bacteria suggest that they can be used to determine the relative dominance of these bacterial groups in sediments (Rajendran *et al.*, 1997). The MUFA/BRFA, MUFA/SSFA, and BRFA/SSFA ratios are approximately 9.3, 1.8, and 0.2 for aerobic bacteria, but these ratios are approximately 0.2, 1.4, and 6.8 for anaerobic bacteria, SRB, and Gram-positive bacteria (Rajendran *et al.*, 1997). In this study, the MUFA/BRFA, MUFA/SSFA, and BRFA/SSFA ratios varied from 0.2 to 1.0, 0.2 to 0.4, and 0.3 to 1.3 for the BFAs, but the BRFA/SSFA ratios ranged from 0.1 to 0.2 for the FFAs (Tab. 2). These results indicate that FAs derived from algae were the dominant components of the aquatic FAs. As previously noted, anoxic conditions in lake sediments may lead to better preservation of some of these algal FAs (Gong and Hollander, 1997), and anaerobic bacteria were the main contributors of FAs from bacterial sources.

Even carbon numbers dominate the FA composition, and CPI₁₆₋₂₈ values from 8.5 to 17.1 and from 8.5 to 17.9 were noted in the BFAs and FFAs, respectively (Tab. 2). The CPI₁₆₋₂₈ values of the BFAs and FFAs were higher than the CPI₂₅₋₃₁ values for the *n*-alkanes. The greater microbial alteration of *n*-alkanes implied by the smaller *n*-alkane CPI values agrees with the larger FA CPI values that signify greater microbial production of secondary acids (Zhou *et al.*, 2010). The MAWT increased gradually as the latitude decreased. A decrease in latitude could lead to greater decomposition of terrigenous OM, as warmer water temperatures would encourage more microbial activity (Zheng *et al.*, 2007). The CPIs summarize the relative proportions of even-numbered and odd-numbered carbon molecules in the FAs and the *n*-alkanol and *n*-alkane distributions of the sedimentary samples. Because the rates of microbial degradation and alteration can differ under non-uniform climate regimes, the CPI values of sediment lipids can reflect relative changes in water temperature. At low water temperatures, microbial degradation and diagenesis of OM are slowed (Kuder and Kruege, 1998), leading to high CPI values. In contrast, low CPI values are evidence of amplified microbial degradation and alteration of OM in lake sediments corresponding to higher water temperatures. The *n*-alkanol CPI₁₆₋₂₈ values

could not be calculated for free *n*-alkanols in some of the sedimentary samples because of the absence of odd-numbered carbon chains. The CPI₁₆₋₂₈ values ranged from 1.9 to 130.9 for the bound *n*-alkanols and from 5.0 to 32.0 for the partially free *n*-alkanols (Tab. 2). Because of the lower MAWT, the *n*-alkanol CPI₁₆₋₂₈ values were higher in the sediments obtained from the northwestern plateau lakes.

The apparent differences in the distributions of *n*-alkanes, *n*-alkanols, and FAs likely reflect a combination of continual microbial degradation of FAs and *n*-alkanols in the sediments and partial replacement of originally deposited FAs and *n*-alkanols by secondary, microbial FAs and *n*-alkanols. In contrast, the *n*-alkane concentrations are less likely to be diminished by degradation and are unlikely to be replaced by microbial contributions (Tenzer *et al.*, 1999). There were marked differences in the results from different biomarkers. The FAs and *n*-alkanols suggest that the OM in most of the sediment samples came mainly from algal and bacterial inputs. However, the *n*-alkane distributions suggest that the OM in the sediments from the northwestern plateau lakes came mostly from terrestrial inputs, while that in the southwestern plateau lakes came mainly from algal and bacterial sources. The differences between the results might be due to the various mechanisms of degradation that affect different biomarkers or a specific input of a certain biomarker; thus, care must be exercised when using a single biomarker as an indicator of OM input.

CONCLUSIONS

The present study addresses the origin and distribution of dissolvable organic matter found in sediment collected from southwestern and northwestern plateau lakes in China at the level of elements, stable carbon isotopes and molecular compositions. The distributions of the components of free and bound lipid biomarkers suggest that the lipids found in the sediment samples are derived from a mixture of terrigenous and aquatic inputs.

The FAs and *n*-alkanols suggest that the OM in most of the sediment samples originated mainly from algal and bacterial inputs, and the aquatic OM fractions were higher in sediments from the southwestern plateau lakes than from the northwestern plateau lakes. This difference may be due to the higher light intensities and MAWTs in the southwestern lakes, which promote the growth and reproduction of algae and bacteria and simultaneously lead to additional decomposition of terrestrial OM. The OM derived from algae is the major component of sedimentary OM, and anaerobic bacteria are the primary contributors of bacterial OM. The compositions of *n*-alkane biomarkers are well preserved in the sediment samples. However, the *n*-alkanols are not as well preserved, and the original, vascular plant *n*-FAs have been largely replaced by mi-

crobial compounds. Components produced by microbial reworking were observed in the bound *n*-alkanols and FAs, but few such components are present in the free *n*-alkanols and FAs.

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